

1 **The rapid formation of macromolecules in irradiated ice of protoplanetary**
2 **disk dust traps**

3
4 **Authors:** Niels F.W. Ligterink^{1,2,*}, Paola Pinilla³, Nienke van der Marel⁴, Jeroen Terwisscha
5 van Scheltinga^{4,5}, Alice S. Booth^{4,6}, Conel M. O'D. Alexander⁷, My E.I. Riebe⁸

6 **Affiliations:**

7 ¹Physics Institute, Space Research and Planetary Sciences, University of Bern, Sidlerstrasse 5,
8 CH-3012 Bern, Switzerland

9 ORCID: 0000-0002-8385-9149

10
11 ²Faculty of Aerospace Engineering, Delft University of Technology, Delft, The Netherlands

12 *email: niels.ligterink@unibe.ch, niels.ligterink@tudelft.nl

13
14 ³Mullard Space Science Laboratory, University College London, Holmbury St Mary, Dorking,
15 Surrey RH5 6NT, UK

16 ORCID: 0000-0001-8764-1780

17
18 ⁴Leiden Observatory, Leiden University, P.O. box 9513, 2300 RA Leiden, The Netherlands

19 ORCID: 0000-0003-2458-9756

20
21 ⁵Laboratory for Astrophysics, Leiden Observatory, Leiden University, PO Box 9513, 2300 RA
22 Leiden, The Netherlands

23 ORCID: 0000-0002-3800-9639

24
25 ⁶Harvard-Smithsonian Center for Astrophysics 60 Garden St, Cambridge, MA 02138, USA

26 ORCID: 0000-0003-2014-2121

27
28 ⁷Earth and Planets Laboratory, Carnegie Institution for Science, 5241 Broad Branch Road NW,
29 Washington, DC 20015, USA

30 ORCID: 0000-0002-8558-1427

31
32 ⁸Institute of Geochemistry and Petrology, Eidgenössische Technische Hochschule Zürich, 8092
33 Zürich, Switzerland

34 ORCID: 0000-0002-2098-9587

35 Abstract

36 **Organic macromolecular matter is the dominant carrier of volatile elements such as carbon,**
37 **nitrogen, and noble gases in chondrites – the rocky building blocks from which Earth**
38 **formed. How this macromolecular substance formed in space is unclear. We show that its**
39 **formation could be associated with the presence of dust traps, which are prominent**
40 **mechanisms for forming planetesimals in planet-forming disks. We demonstrate the**
41 **existence of heavily irradiated zones in dust traps, where small frozen molecules that coat**
42 **large quantities of microscopic dust grains could be rapidly converted into macromolecular**
43 **matter by receiving radiation doses of up to several 10s of eV molecule⁻¹ year⁻¹. This allows**
44 **for the transformation of simple molecules into complex macromolecular matter within**
45 **several decades. Approximately and up to 4% of the total disk ice reservoir can be processed**
46 **and inherited into the protoplanetary disk midplane in this way. This finding shows that**
47 **planetesimal formation and the production of organic macromolecular matter, which**
48 **provides essential elemental building blocks of life, might be linked.**

49 Introduction

50 Organic macromolecular matter likely supplied the terrestrial planets with most of their carbon,
51 nitrogen, and noble gases¹. In chondrites, this material is often **called** Insoluble Organic Matter
52 (IOM), where it is the dominant carrier of these volatile elements². It has even been suggested that
53 organic macromolecular matter directly contributed to the emergence of life³. Similarities in the
54 elemental compositions of chondritic IOM and refractory organic matter in comets, as well as large
55 deuterium enrichments, indicate that there is a genetic relationship between the two materials⁴. A
56 genetic relationship is, perhaps, not so surprising since the formation regions of some chondrites
57 were probably well beyond the orbit of Jupiter⁵. Hence, refractory organic matter was distributed
58 over large radial distances in the proto-Solar Nebula⁶. Detailed characterization of IOM has
59 provided constraints for potential formation mechanisms, but there is still no consensus about
60 whether this material formed in the interstellar medium, the proto-Solar Nebula, or the parent
61 bodies by polymerization of simpler precursors². **With a model, we demonstrate that**
62 **macromolecular organic matter resembling that of IOM in chondrites² and refractory**
63 **organics in comets⁶ could rapidly form in dust traps in the proto-Solar Nebula, via radiation-**
64 **driven ice chemistry. Large quantities of ice, on the order of and up to 4% of the total disk**
65 **ice reservoir in our model, could be processed and inherited in this way. Our findings suggest**
66 **a link between the mechanism that forms planetesimals and the chemical processes that set**
67 **their macromolecular and volatile element budget.**

68 **Due to the steady influx of meteorites to Earth,** IOM is the best-studied of all macromolecular
69 organic matter. Its structure is characterized by small (poly)cyclic aromatic units, linked by short,
70 highly branched aliphatic chains, furan/ether bonds, and additional functional groups, such as
71 ketones and carboxyls^{7,8}. Deuterium and ¹⁵N enhancements in bulk, and even larger ones in
72 individual grains, require a cold (<20 K) environment where isotopes are readily fractionated, and
73 simple molecules are frozen onto dust grains⁹. Subsequent radiation of frozen organics could
74 enhance deuterium enrichments¹⁰, while heating and aqueous alteration in the parent bodies may
75 have reduced them¹¹. IOM contains high concentrations of radicals¹² that may have been generated
76 during irradiation of organic molecules¹³. IOM also contains high concentrations¹³ of noble gases
77 which are isotopically fractionated and depleted in light noble gases compared to the Solar

78 composition¹⁴, which is suggestive of a thermal loss mechanism starting at low temperatures (<100
79 K).

80 These observations suggest that the formation of IOM and refractory organics in comets, involves
81 the irradiation of frozen carbon-bearing molecules. **Laboratory studies of particle and UV**
82 **irradiation of ice films and frozen organic molecules have demonstrated that these molecules**
83 **can be converted into macromolecular matter¹⁵⁻¹⁸, although mixing ratios are not always**
84 **realistic with respect to the ice found in interstellar environments and protoplanetary disks.**
85 **Radiation doses of several 100s up to 1000 eV molecule⁻¹ are required for the conversion,**
86 **which are much higher than what molecules in the solid phase typically experience during**
87 **the star- and planet-formation cycle.** UV and cosmic-ray doses received by ice-coated grains in
88 dark clouds and protoplanetary disk midplanes during their lifetimes at most reach several 10s of
89 eV molecule⁻¹ ($\sim 10^{-7}$ eV molecule⁻¹ year⁻¹)¹⁹. Radiation doses received by ice-coated dust grains
90 in the outer layers of a protoplanetary disk, which are directly irradiated by the protostar, **are**
91 **similarly moderate and only sufficient to produce much smaller and typically solvent-soluble**
92 **organic molecules (SOM, such as amino acids, sugars, polyoxymethylene, and other**
93 **polymers)²⁰.** Until now, it was unknown where large quantities of ice and other frozen materials
94 could be heavily irradiated to form organic macromolecular matter during the star- and planet-
95 formation sequence.

96
97 **Following the first detection of a dust trap in the protoplanetary disk IRS48²¹, our**
98 **understanding of planetesimal formation has been revolutionized²².** Dust traps are localized
99 pressure bumps in protoplanetary disks where the radial drift of dust is reduced or stopped, and
100 material piles up²³. Grains can efficiently grow to form pebbles and planetesimals^{22,24}. Dust traps
101 are regularly found in protoplanetary disks with different properties (e.g., disk ages from < 1 to 10
102 Myr) and in the form of rings and crescents in millimeter dust continuum observations with
103 ALMA²². **They** are thought to have played a fundamental role in the formation of the Solar System,
104 as they were likely present in the primordial disk, even beyond the current orbit of Uranus²⁵. Based
105 on their distinct isotopic compositions, the different chondrite groups that host IOM have been
106 suggested to have collected in a dust trap that formed in the first 2-4 Myr²⁶.

107
108 Dust traps have been observationally shown to contain large amounts of ice-associated molecules
109 that were likely inherited from the molecular cloud stage²⁷. H₂CO, CH₃OH, and CH₃OCH₃ have
110 recently been detected in the gas-phase in a region co-spatial with dust traps²⁸⁻³⁰. **Their detection**
111 **can be explained by the vertical transport of ice-coated grains and thermal desorption in the**
112 **warmer surface layers, whereas in disks without dust traps these molecules remain locked**
113 **up in ice-coated grains, thus remaining undetectable with ALMA³¹.** Irradiated dust traps could
114 thus provide favourable conditions for the formation of organic macromolecules.

115 In this study, we use the output of a state-of-the-art dust evolution model of a protoplanetary disk
116 with a dust trap³² and calculate the UV radiation dose rate in ice-coated grains throughout the disk.
117 **While we knew of radiation-heavy environments before and ice-rich environments, there**
118 **was no known environment where high radiation doses and large ice reservoirs came**
119 **together. We demonstrate the existence of heavily irradiated regions in dust traps, where**
120 **large ice reservoirs can rapidly be transformed into organic macromolecular matter such as**
121 **IOM.**

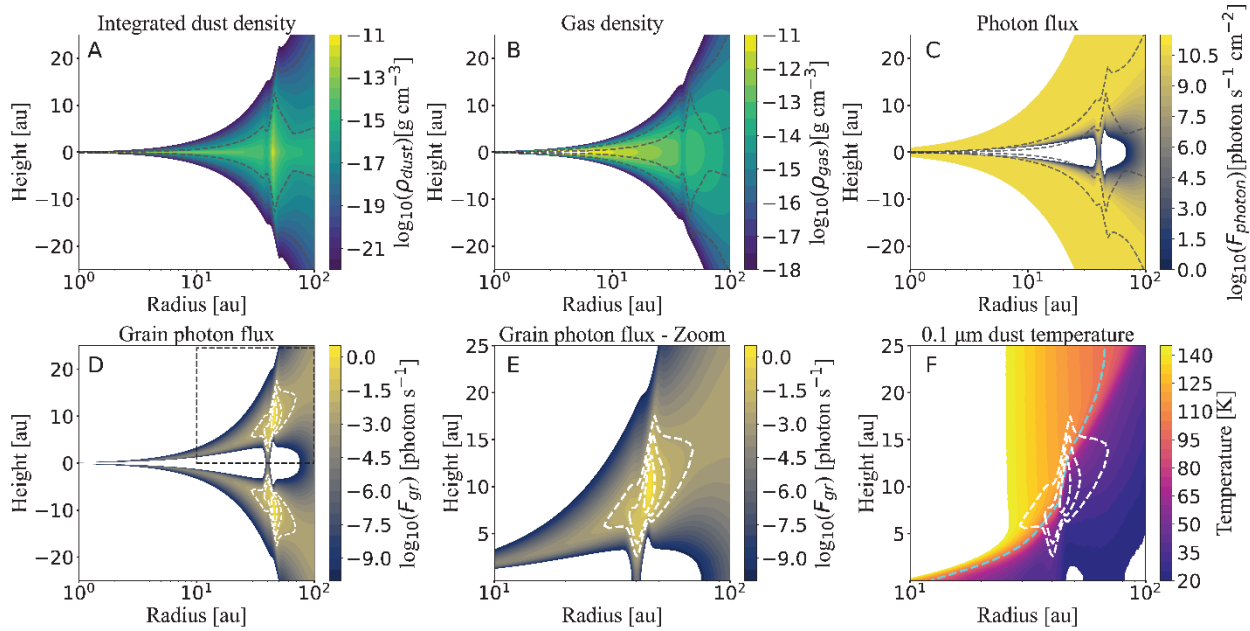
122

123 Results

124 The dust and gas distributions predicted by the dust evolution model for a 1 Myr old protoplanetary
125 disk with a dust trap at 45 au are shown in **Fig. 1.A** and **1.B** (details in the **Methods**). **The dust-**
126 **to-gas ratio exceeds unity in the dust trap of this model, which is a sign that planetesimal**
127 **formation ensues.** The modelled dust grain sizes range from 10^{-7} to 10^{-1} m, but grains smaller
128 than several tens of μm dominate at elevated heights in the dust trap (**Fig. S1**). The assumption is
129 made that each dust grain is covered by 100 monolayers (ML, $1 \text{ ML} = 10^{15} \text{ molecules cm}^{-2}$) of
130 frozen molecules. **Since we determine the average dose rate per molecule, the composition of**
131 **the ice is arbitrary, but it can be thought to resemble those of interstellar dust grains,** that is,
132 H_2O dominated, with fractions of CH_4 , CO , CO_2 , CH_3OH , and NH_3 . Alternatively, **warm** grains
133 can be coated in a layer of non-volatile organic molecules after water ice has sublimated³³. UV
134 photons emitted by the protostar penetrate throughout the disk (**Fig. 1.C**) and the photon flux is
135 calculated as a vertical radiation field, which is attenuated by the gas and dust throughout the disk.
136 The optical depth is determined following $\tau(\mathbf{r}, \mathbf{z}) = \int_z^\infty \rho(\mathbf{r}, \mathbf{z}) \kappa d\mathbf{z}$ where the column of
137 material at a radius R from the central star above a height Z above the midplane is multiplied with
138 the average opacity κ ($\text{cm}^2 \text{ g}^{-1}$). The UV flux is calculated with the Beer-Lambert law $\mathbf{F}(\mathbf{r}, \mathbf{z}) =$
139 $\mathbf{F}_0 e^{-\tau(\mathbf{r}, \mathbf{z})}$, where F_0 is the number of incoming photons. In the fiducial model, we use $F_0 = 1000$
140 G_0 , where G_0 is a UV flux of $10^8 \text{ photons cm}^{-2} \text{ s}^{-1}$. Various models of protoplanetary disks
141 demonstrate that this is a flux that is readily achieved in the surface layers of disks, due to the
142 radiation from the central star³⁴. The mean opacity is set to $\kappa = 10 \text{ cm}^2 \text{ g}^{-1}$, which is a value
143 commonly used for protoplanetary disks³⁵.

144 The grain photon flux (product of the photon flux and the grain surface areas) of the fiducial model
145 is shown in **Figs. 1.D,E**. It reveals that the largest reservoir of ice-coated dust grains receiving the
146 largest photon flux is in the dust trap at $Z = \pm 10$ astronomical units (au) and $R = 45$ au. This is the
147 result of local gas densities being lower near the inner edge of the dust ring or dust trap, while dust
148 densities are large. In comparison, gas densities in the inner disk (taken as $r < 10$ au in our model)
149 are large (see **Fig. 1.B**) and minor quantities of disk ice receive an appreciable photon flux (see
150 **Fig. 1.D**). In the dust trap at a height of $Z = 5 - 20$ au, grain temperatures generally exceed ~ 50
151 K, while sub- μm grain temperatures surpass even ~ 100 K in specific regions (see **Fig. 1.F** and
152 **Fig. S2**). Therefore, most grains in the dust trap are ‘luke-warm’ **and retain ice films that are**
153 **dominated by less volatile species (e.g., H_2O , CH_3OH , PAHs), small amounts of trapped more**
154 **volatile species, and minor quantities of other organic molecules**³³. **The combination of high**
155 **dust concentrations, moderately elevated temperatures, and radiation makes the dust trap**

156 wall a hotspot for radiation-driven ice chemistry, which stimulates the formation of complex
 157 molecules³⁶.

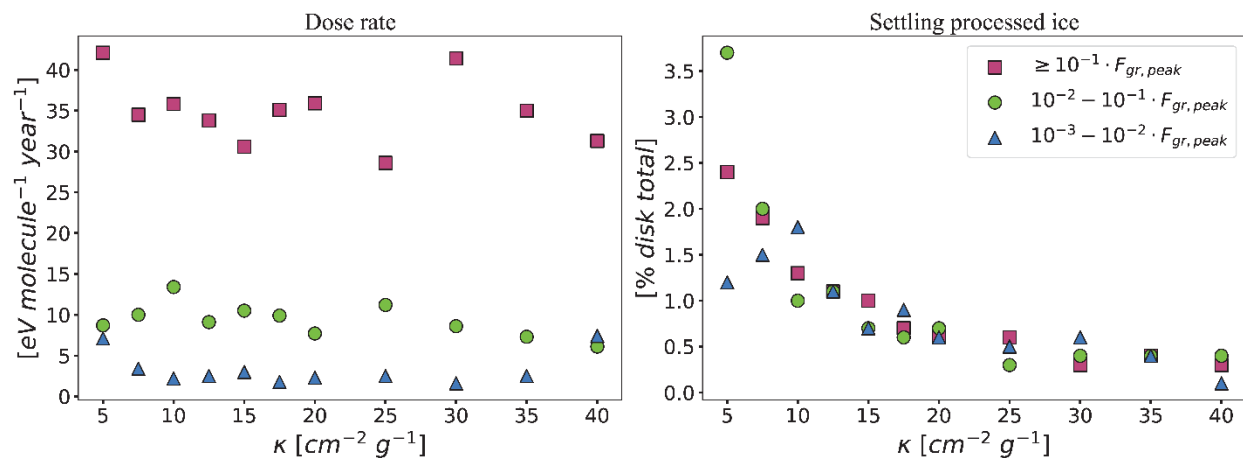


158
 159 **Fig. 1.** Dust evolution and irradiation model of a protoplanetary disk with a dust trap located at
 160 ~ 45 au. The integrated dust density distribution (A), gas density (B), photon flux for an opacity of
 161 $\kappa = 10 \text{ cm}^2 \text{ g}^{-1}$ (C), grain photon flux (D,E), and temperature of the $1 \mu\text{m}$ dust grains (F) are
 162 shown. The black contours indicate the dust density at 10^{-19} (outer) and 10^{-16} (inner) g cm^{-3} , the
 163 **hotspot** region is depicted by white contours at 10^{-3} , 10^{-2} , and 10^{-1} times the peak grain photon
 164 flux, and the cyan contour indicates the boundary where the grain temperature equals 100 K.

165 The dust trap **hotspot** is divided into regions at intervals of 10^{-3} , 10^{-2} , and 10^{-1} times the peak
 166 grain flux ($F_{\text{gr,peak}}$). By dividing the incoming photon flux by the amount of ice and assuming an
 167 average UV photon energy of 6 eV ³⁷, the average energetic input is determined in eV molecule^{-1}
 168 s^{-1} for all molecules in the ice mantle, including H_2O . At its peak position, the fiducial model shows
 169 that the dose rate is $36 \text{ eV molecule}^{-1} \text{ yr}^{-1}$. A dose of $1000 \text{ eV molecule}^{-1}$, which is sufficient to
 170 produce macromolecular matter¹⁵⁻¹⁷, is achieved within 30 years. Even in the extended region,
 171 where the dose rate decreases to $2 \text{ eV molecule}^{-1} \text{ yr}^{-1}$, **this** dose is obtained in approximately 500
 172 years. These timescales are well within the expected lifetime of a dust trap²³.

173 A substantial quantity of ice is contained in the **hotspot** region, with 2.6%, 2.1%, and 3.5% of the
 174 total disk ice reservoir in the three regions, respectively (8% total). Grains are subject to vertical
 175 stirring, which results in the loss of ice-coated grains to the disk atmosphere when grains are stirred
 176 upwards. For this model, approximately half of the grains are lost (see **Methods** for details), **which**
 177 **means that the other half (4%) of the processed ice reservoir settles to the disk midplane,**
 178 where it is available as planetesimal-forming material. Because such a large quantity of ice is
 179 processed, organic macromolecular matter formation in dust traps can currently best explain the
 180 $\sim 6\%$ formation efficiency of IOM and up to 55% for comets as derived from atomic C/Si ratios².
 181 The ice-conversion factor is likely higher in a dynamic disk where fresh ice-coated grains are
 182 continuously replenished in the dust trap **hotspot** due to vertical and radial mixing. **Conversion**
 183 **efficiencies of pristine ice to organic macromolecular matter have not been determined and**

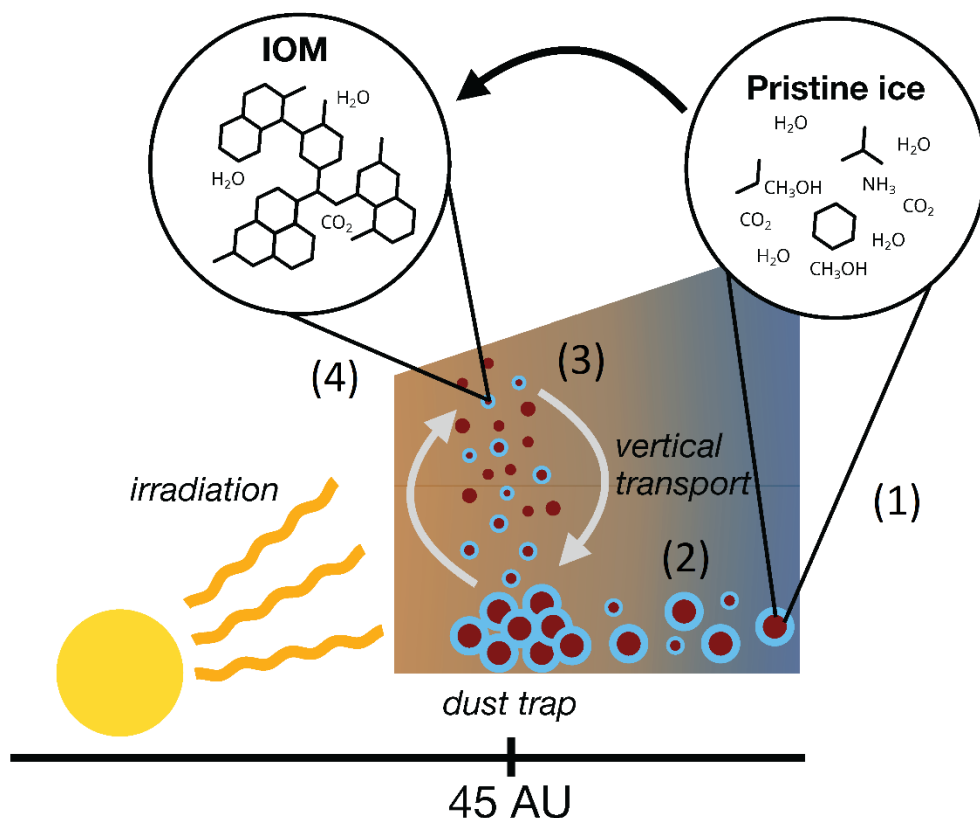
184 **therefore the amount of processed ice is an upper limit of produced macromolecular matter.**
 185 Furthermore, dust trap locations and opacities may differ from disk to disk and will critically
 186 influence the dose rates and above-derived conversion fraction. For opacities ranging from 5 – 40
 187 $\text{cm}^2 \text{g}^{-1}$, the average dose rate does not alter significantly for this model (**Fig. 2.A.**), but the amount
 188 of processed frozen material increases with decreasing κ , see **Fig. 2.B.** This is the result of photons
 189 penetrating deeper into the disk, that is, closer to the midplane, and accessing regions where the
 190 concentration of dust grains is larger.



191
 192 **Fig. 2.** Influence of the opacity on the dose rate and amount of settled processed ice for regions
 193 constrained between 10^{-3} , 10^{-2} , and 10^{-1} the peak grain photon flux. While the dose rate remains

194 constant, the amount of processed ice that can settle back to the midplane increases with decreasing
195 disk opacity.

196



197

198 **Fig. 3.** Schematic depiction of the IOM formation scenario. Grains with pristine, simple ice
199 components (1) radially drift into the protoplanetary disk (2) and migrate to the dust trap **hotspot**
200 through vertical mixing/transport (3), where heavy irradiation of the luke-warm ice results in the
201 formation of organic macromolecular matter (4).

202 Discussion

203 The following scenario for the formation of macromolecular matter is proposed (**Fig. 3.**). In dense
204 interstellar clouds, dust grains acquire ice mantles of simple species that are enriched in minor
205 isotopes^{38,39}. The ice or dust grain may be enriched in PAHs. Noble gases are efficiently trapped
206 in these ice mantles⁴⁰ and moderate grain surface chemistry results in the production of minor
207 quantities of SOM⁴¹. The protoplanetary disk inherits these ice-coated grains from its host cloud²⁷.
208 As grains migrate into the dust trap **hotspot**, the ice warms up and the UV flux increases, **resulting**
209 **in a rapid conversion of the ice mantle material into macromolecular matter, by producing**
210 **and linking complex ice-irradiation products, such as low volatility SOM, and PAH**
211 **fragments.** Noble gases in the ice mantles may be **trapped** in the macromolecular matter during

212 this process⁴², and D and ¹⁵N isotopic signatures could be fractionated during further irradiation
213 due to preferential loss of light isotopes^{10,13}. Processed grains are cycled back into the midplane,
214 creating a mixture of pristine ice and macromolecular matter, in line with the fact that chondrites
215 accreted significant amounts of ice⁴³. Migration of these mixtures closer to the protostar can result
216 in IOM-rich objects like asteroid Ryugu⁴⁴, whereas staying at these large radii (45 au in this model)
217 preserves it as refractory organics as seen in comets⁴⁵. Since multiple dust traps can occur in a
218 protoplanetary disk, as well as in the proto-Solar Nebula (inside or outside the water ice line),
219 cometary refractory organics and IOM in meteorites and asteroids may be formed at different radii
220 and times. In both cases, the heavy irradiation of grain mantles results in the production of similar
221 macromolecules.

222 **IOM shows a range** of chemical characteristics^{2,46}, which can be explained by this scenario. Disks
223 are turbulent environments²³, **creating** variations in photon flux and grain surface temperature.
224 Protoplanetary disks can have multiple dust traps at different radii with different shapes and
225 amplitudes, and the flux of radiation from the central star can vary with time²², which would further
226 enhance the diversity of macromolecular products and their characteristics. For example,
227 deuterium fractionation by irradiation of organic polymers has been demonstrated, where different
228 radiation fluxes will result in different levels of fractionation¹⁰. The mixing of material that
229 experienced different degrees of alteration of these pathways explains the heterogeneous isotope
230 nature of IOM on (sub)µm scales⁴⁶. Small (sub)µm ice-coated dust grains are preferentially stirred
231 up into the **hotspot** region [Fig. S1.], **but trajectories and residence times may vary from a few**
232 **to several dozen Ω_k^{-1} (in this model, Ω_k^{-1} is ~30-50 years at the hotspot radius)⁴⁷. However,**
233 **because small grains are continuously re-created in the trap due to the fragmentation of the**
234 **large particles, there is a continuous amount of small grains during million-year timescales**
235 **that are exposed to the irradiation needed for IOM formation.** This affects the transformation
236 of pristine material into organic macromolecular matter and, in turn, its chemical composition and
237 the texture of the material. For example, nanoglobules, small spherical and sometimes hollow
238 inclusions^{48,49}, have been suggested to form by UV irradiation of ice mantles of (sub)micrometer-
239 sized grains⁴⁹, although **alternative** formation pathways by aqueous alteration have been
240 suggested as well⁸.

241 A prominent question that remains, is whether organic macromolecular matter is formed in a
242 water-rich or -poor environment. Ice mixtures in irradiation experiments that produce IOM
243 analogues contain minor **or no fractions** of water^{16,17} **and do not match with the known ice**
244 **composition of ice-coated dust grains or comets³³.** Water-rich ice is expected to lead to
245 **enhanced CO₂ formation and** oxygen-rich macromolecular material, inconsistent with **the**
246 **carbon-rich elemental composition of organic macromolecular matter observed in space.**
247 **From the experimental work conducted thus far, it is unknown whether macromolecular**
248 **matter can form in water-dominated ice. Instead, its** formation may rely on a preceding step
249 where low volatility SOM is formed, water-ice is removed (e.g., by thermal desorption)⁴¹, and
250 followed by heavy irradiation to form complex refractory macromolecules^{15,18}. In our model, some
251 of the smallest dust grains, which dominate the dust composition in the **hotspot**, reach
252 temperatures at which solid-state water readily sublimates. If a dust trap is located closer to the
253 protostar, grains of larger sizes will be heated up to higher temperatures and the loss of water-ice
254 becomes more prominent, **while organics of low volatility remain.** A thermal heating cycle also
255 **helps** explain why noble gases detected in natural IOM are depleted in the lighter species compared

256 to the Solar composition since He and Ne desorb at significantly lower temperatures compared to
257 Kr and Xe.

258 **To determine the required dose to form organic macromolecular matter, we cite particle**
259 **irradiation studies¹⁵⁻¹⁷, whereas our model relies on UV radiation. Photons are affected by**
260 **the optical properties of the ice, whereas penetration depths of particles of keV and larger**
261 **energies are usually larger than UV photons. These differences may result in different**
262 **chemical outcomes of ice irradiation. However, comparative experiments products are the**
263 **same for ice films irradiated with UV photon or ions at similar dose⁵⁰. This makes it plausible**
264 **that both radiation types will result in a similar type of organic macromolecular matter, as**
265 **indirectly indicated by the UV irradiation experiments¹⁸.**

266 Over the last years, high-resolution observations of protoplanetary disks have demonstrated that
267 disks have dust traps across different disk and stellar properties. **We demonstrate that hotspots,**
268 **that is, zones of heavy irradiation and increased dust temperatures, exist in these dust traps,**
269 **where the icy mantle material on dust grains can rapidly be converted to organic**
270 **macromolecular matter. This suggests that the mechanism by which planets form also**
271 **produces the macromolecular matter from which terrestrial planets derive their elemental**
272 **carbon and nitrogen, which contribute to the emergence of life.**

273 **Methods**

274 The dust evolution models are performed using the Dustpy code⁵¹. The disk is assumed to be
275 around a Solar-mass star, with a gas surface density that assumes a critical radius at 80 AU and an
276 initial disk mass of 5% of the Sun²³. The model assumes a fragmentation velocity of $v_f = 10 \text{ m s}^{-1}$,
277 a gas viscous evolution parameter, radial diffusion, turbulent mixing, and vertical settling/stirring
278 all set to 10^{-4} . All grains are initially small between 0.1-1 microns in size, with a power law
279 distribution as $n(a) \propto a^{-3.5}$. The models include the dust growth and dynamics of particles. The
280 model output gives the density distribution of the dust grains (ρ_{dust} , **Fig. 1.A**) and of the gas (ρ_{gas} ,
281 **Fig. 1.B**) at 1 Myr of evolution. The data is **portrayed** in a radial grid from 1 to 300 au (only 1 to
282 100 au shown in **Fig. 1**) with 300 logarithmically spaced cells. A logarithmically spaced grain size
283 distribution between the minimum grain size of **0.1 micron to 0.1 m** in 127 steps is used. A gap
284 **centred** at 40 AU is assumed in the disk³², which yields a pressure bump at around 45 AU. Smaller
285 and lighter dust grains are more easily stirred up in the disk and therefore dominate the particle
286 sizes at greater disk heights, whereas larger particles concentrate around the disk midplane [**Fig.**
287 **S1.**].

288
289 From the dust density distribution, the grain surface area is calculated by assuming that all grains
290 are spherical and have a density of $\rho_{\text{grain}} = 1.65 \text{ g cm}^{-3}$. This allows us to calculate the particle mass
291 for each grain size and convert the dust density (g cm^{-3}) into a particle density (n cm^{-3}). Next, the
292 particle density is multiplied by the grain area to find the total grain area per volume, which gives:

$$295 \quad A_{\text{grain}} = \frac{(3 \cdot \rho_{\text{dust}})}{(\rho_{\text{grain}} \cdot r_{\text{grain}})} \quad (1)$$

296

297 where r_{grain} is the grain size (that is, the radius). Using the grain area, the available amount of ice
 298 is calculated by assuming that each grain is covered with 100 monolayers (1 ML = 10^{15} molecules
 299 cm^{-2}) of ice and by multiplying this value with the available grain surface area.

300
 301 The thermal structure of the disk is modelled with RADMC-3D⁵², assuming a vertical grid of 180
 302 cells over a semicircle following $Z = R \cdot \cos(\theta)$, where Z is the disk height and R the radius, and
 303 following the same procedure as³². The grain sizes that dominate the dust trap **hotspot** have sizes
 304 of 0.1 to several tens of μm and generally have temperatures greater than 50 K. However, only
 305 sub- μm grains are heated above 100 K, and only in specific regions of the **hotspot** [Fig. S2.]
 306

307 The **number** of photons throughout the disk is calculated with the Beer-Lambert law (see **main**
 308 **text**). The impinging photon flux is fixed to $F_0 = 1000 G_0$ ($G_0 = 10^8 \text{ photons cm}^{-2} \text{ s}^{-1}$), but we note
 309 the (grain) photon flux throughout the disk scales linearly with the value chosen for F_0 . Therefore,
 310 the dose rate can be scaled in the same way. The number of photons absorbed by the mantle
 311 depends on its thickness and the UV photon absorption cross-section. Assuming the mantle
 312 consists entirely of water ice, then a film of 100 ML does not fully absorb the received flux⁵³. The
 313 simplified division of the photon energy input by the column density to yield $\text{eV molecule}^{-1} \text{ s}^{-1}$ is
 314 therefore a rough assumption. However, since the photons that penetrate the ice mantle still hit the
 315 underlying grain, we assume that the energy is contributed to the overall system and therefore the
 316 simplified division holds.

317
 318 Gravitational attraction causes grains to settle in the disk midplane, while turbulent stirring can
 319 move particles towards or away from the midplane. The velocities of these processes can be
 320 calculated⁵⁴ and in turn, be used to assess how the dust is vertically distributed.

321
 322 **The equation to determine the velocity with which dust settles to midplane is given as:**

$$323 \quad v_{\text{sett}} = z \Omega_{\text{k}} ST, \quad (2)$$

324
 325 where z is the vertical height in meter and Ω_{k} is the Keplerian frequency:

$$326 \quad \Omega_{\text{k}} = \sqrt{\frac{G M_{\odot}}{R^3}}, \quad (3)$$

327
 328 with G the gravitational constant, M_{\odot} the Solar mass in kg, and R the radius in m, and ST
 329 the Stokes parameter, calculated following:

$$330 \quad ST = \frac{r_{\text{grain}} \rho_{\text{grain}} \pi}{\Sigma_g} \frac{\pi}{2} \quad (4)$$

331
 332 where Σ_g the gas surface density.

333
 334 To calculate the vertical stirring velocity, we assume that this velocity is as the relative
 335 velocities due to turbulence of particles of similar size, and use the following equation:

336
 337
 338
 339
 340

341

$$v_{stir} = \sqrt{\frac{3\alpha}{ST+ST^{-1}}} c_S \quad (5)$$

342

343 is used, where α is the gas viscous evolution parameter ($\alpha = 10^{-4}$) and c_S is the isothermal
344 sound velocity:

345

346

347

$$c_S = \sqrt{\frac{k_B T_{gas}}{\mu_{gas} m_{proton}}}, \quad (6)$$

348

349 with T_{gas} the gas temperature ($T_{gas} = 50$ K, the average dust temperature at the trap
350 location and assuming $T_{gas} = T_{dust}$), μ_{gas} the mean molecular mass of the gas ($\mu_{gas} = 2$), and
351 m_{proton} the proton mass in kg. Note that in the disk surface, where our calculations are
352 relevant, the dust densities are dominated by the small grains, so the dust particles have
353 similar Stokes numbers, which endorse the use of Eq (5) for the stirring velocities.

354

355 For the model results used in this work, $v_{stir} \gg v_{sett}$ for the μm -sized particles that reside at greater
356 height in the dust trap. Therefore, stirring determines which direction material migrates. Since
357 vertical stirring can point into two directions, namely away from the midplane (up) or towards the
358 midplane (down), we assume that the likelihood of up- or downward motion is equal. This means
359 that 50% of material in the dust trap **hotspot** goes to the disk atmosphere and 50% moves towards
360 the midplane. We note that this situation holds for a static snapshot of the disk model, but in a
361 dynamic and evolving environment the loss and settle fractions may be different.

362

363

364 References and Notes

- 365 1. Marty, B. The origins and concentrations of water, carbon, nitrogen and noble gases on
366 Earth. *Earth Planet Sci Lett* **313**, 56–66 (2012).
- 367 2. Alexander, C. M. O., Cody, G. D., De Gregorio, B. T., Nittler, L. R. & Stroud, R. M. The
368 nature, origin and modification of insoluble organic matter in chondrites, the major source
369 of Earth's C and N. *Geochemistry* **77**, 227–256 (2017).
- 370 3. d'Ischia, M. *et al.* Insoluble organic matter in chondrites: Archetypal melanin-like PAH-
371 based multifunctionality at the origin of life? *Phys Life Rev* **37**, 65–93 (2021).
- 372 4. Paquette, J. A. *et al.* D/H in the refractory organics of comet 67P/Churyumov-
373 Gerasimenko measured by Rosetta/COSIMA. *Mon Not R Astron Soc* **504**, 4940–4951
374 (2021).
- 375 5. Walsh, K. J., Morbidelli, A., Raymond, S. N., O'Brien, D. P. & Mandell, A. M. A low
376 mass for Mars from Jupiter's early gas-driven migration. *Nature* **475**, 206–209 (2011).
- 377 6. Kissel, J. & Krueger, Fr. The organic component in dust from comet Halley as measured
378 by the PUMA mass spectrometer on board Vega 1. *Nature* **326**, 755–760 (1987).
- 379 7. Hayatsu, R., Matsuoka, S., Scott, R. G., Studier, M. H. & Anders, E. Origin of organic
380 matter in the early solar system—VII. The organic polymer in carbonaceous chondrites.
381 *Geochim Cosmochim Acta* **41**, 1325–1339 (1977).

- 382 8. Cody, G. D. *et al.* Establishing a molecular relationship between chondritic and cometary
383 organic solids. *Proceedings of the National Academy of Sciences* **108**, 19171–19176
384 (2011).
- 385 9. Robert, F. & Epstein, S. The concentration and isotopic composition of hydrogen, carbon
386 and nitrogen in carbonaceous meteorites. *Geochim Cosmochim Acta* **46**, 81–95 (1982).
- 387 10. Laurent, B. *et al.* The deuterium/hydrogen distribution in chondritic organic matter attests
388 to early ionizing irradiation. *Nat Commun* **6**, 8567 (2015).
- 389 11. Alexander, C., Boss, A. P., Keller, L. P., Nuth, J. A. & Weinberger, A. Astronomical and
390 meteoritic evidence for the nature of interstellar dust and its processing in protoplanetary
391 disks. *Protostars and planets V* 801–813 (2007).
- 392 12. Binet, L., Gourier, D., Derenne, S., Pizzarello, S. & Becker, L. Diradicaloids in the
393 insoluble organic matter from the Tagish Lake meteorite: Comparison with the Orgueil
394 and Murchison meteorites. *Meteorit Planet Sci* **39**, 1649–1654 (2004).
- 395 13. Laurent, B. *et al.* Isotopic and structural signature of experimentally irradiated organic
396 matter. *Geochim Cosmochim Acta* **142**, 522–534 (2014).
- 397 14. Ott, U. Planetary and pre-solar noble gases in meteorites. *Geochemistry* **74**, 519–544
398 (2014).
- 399 15. Strazzulla, G. & Baratta, G. A. Carbonaceous material by ion irradiation in space.
400 *Astronomy and Astrophysics (ISSN 0004-6361)*, vol. 266, no. 1, p. 434–438. **266**, 434–438
401 (1992).
- 402 16. Ferini, G., Baratta, G. A. & Palumbo, M. E. A Raman study of ion irradiated icy mixtures.
403 *Astron Astrophys* **414**, 757–766 (2004).
- 404 17. Palumbo, M. E., Ferini, G. & Baratta, G. A. Infrared and Raman spectroscopies of
405 refractory residues left over after ion irradiation of nitrogen-bearing icy mixtures.
406 *Advances in Space Research* **33**, 49–56 (2004).
- 407 18. Danger, G. *et al.* The transition from soluble to insoluble organic matter in interstellar ice
408 analogs and meteorites. *Astron Astrophys* **667**, A120 (2022).
- 409 19. Shen, C. J., Greenberg, J. M., Schutte, W. A. & Van Dishoeck, E. F. Cosmic ray induced
410 explosive chemical desorption in dense clouds. *Astron Astrophys* **415**, 203–215 (2004).
- 411 20. Ciesla, F. J. & Sandford, S. A. Organic synthesis via irradiation and warming of ice grains
412 in the solar nebula. *Science (1979)* **336**, 452–454 (2012).
- 413 21. van der Marel, N. *et al.* A major asymmetric dust trap in a transition disk. *Science (1979)*
414 **340**, 1199–1202 (2013).
- 415 22. Andrews, S. M. Observations of protoplanetary disk structures. *Annu Rev Astron*
416 *Astrophys* **58**, 483–528 (2020).
- 417 23. Pinilla, P. *et al.* Trapping dust particles in the outer regions of protoplanetary disks. *Astron*
418 *Astrophys* **538**, A114 (2012).
- 419 24. Youdin, A. N. & Goodman, J. Streaming instabilities in protoplanetary disks. *Astrophys J*
420 **620**, 459 (2005).
- 421 25. Izidoro, A. *et al.* Planetesimal rings as the cause of the Solar System’s planetary
422 architecture. *Nat Astron* **6**, 357–366 (2022).
- 423 26. Hellmann, J. L. *et al.* Origin of isotopic diversity among carbonaceous chondrites.
424 *Astrophys J Lett* **946**, L34 (2023).
- 425 27. Booth, A. S. *et al.* An inherited complex organic molecule reservoir in a warm planet-
426 hosting disk. *Nat Astron* **5**, 684–690 (2021).

- 427 28. van Der Marel, N., Booth, A. S., Leemker, M., van Dishoeck, E. F. & Ohashi, S. A major
428 asymmetric ice trap in a planet-forming disk-I. Formaldehyde and methanol. *Astron*
429 *Astrophys* **651**, L5 (2021).
- 430 29. Brunken, N. G. C. *et al.* A major asymmetric ice trap in a planet-forming disk-III. First
431 detection of dimethyl ether. *Astron Astrophys* **659**, A29 (2022).
- 432 30. Booth, A. S. *et al.* Sulphur monoxide emission tracing an embedded planet in the HD
433 100546 protoplanetary disk. *Astron Astrophys* **669**, A53 (2023).
- 434 31. van der Marel, N. Transition disks: the observational revolution from SEDs to imaging.
435 *The European Physical Journal Plus* **138**, 225 (2023).
- 436 32. Pinilla, P., Lenz, C. T. & Stammer, S. M. Growing and trapping pebbles with fragile
437 collisions of particles in protoplanetary disks. *Astron Astrophys* **645**, A70 (2021).
- 438 33. Boogert, A. C. A., Gerakines, P. A. & Whittet, D. C. B. Observations of the icy universe.
439 *Annu Rev Astron Astrophys* **53**, 541–581 (2015).
- 440 34. Bruderer, S. Survival of molecular gas in cavities of transition disks-I. CO. *Astron*
441 *Astrophys* **559**, A46 (2013).
- 442 35. Cevallos Soto, A., Tan, J. C., Hu, X., Hsu, C.-J. & Walsh, C. Inside-out planet formation–
443 VII. Astrochemical models of protoplanetary discs and implications for planetary
444 compositions. *Mon Not R Astron Soc* **517**, 2285–2308 (2022).
- 445 36. Potapov, A., Fulvio, D., Krasnokutski, S., Jäger, C. & Henning, T. Formation of complex
446 organic and prebiotic molecules in H₂O: NH₃: CO₂ ices at temperatures relevant to hot
447 cores, protostellar envelopes, and planet-forming disks. *J Phys Chem A* **126**, 1627–1639
448 (2022).
- 449 37. Ligterink, N. F. W. *et al.* Controlling the emission profile of an H₂ discharge lamp to
450 simulate interstellar radiation fields. *Astron Astrophys* **584**, A56 (2015).
- 451 38. Bacmann, A. *et al.* CO depletion and deuterium fractionation in prestellar cores. *Astrophys*
452 *J* **585**, L55 (2003).
- 453 39. Spezzano, S., Caselli, P., Sipilä, O. & Bizzocchi, L. Nitrogen fractionation towards a pre-
454 stellar core traces isotope-selective photodissociation. *Astron Astrophys* **664**, L2 (2022).
- 455 40. Almayrac, M. G. *et al.* The EXCITING Experiment Exploring the Behavior of Nitrogen
456 and Noble Gases in Interstellar Ice Analogs. *Planet Sci J* **3**, 252 (2022).
- 457 41. Sandford, S. A., Nuevo, M., Bera, P. P. & Lee, T. J. Prebiotic astrochemistry and the
458 formation of molecules of astrobiological interest in interstellar clouds and protostellar
459 disks. *Chem Rev* **120**, 4616–4659 (2020).
- 460 42. Sandford, S. A., Bernstein, M. P. & Swindle, T. D. The trapping of noble gases by the
461 irradiation and warming of interstellar ice analogs. *Meteoritics & Planetary Science*, vol.
462 **33**, p. A135 **33**, A135 (1998).
- 463 43. Sridhar, S., Bryson, J. F. J., King, A. J. & Harrison, R. J. Constraints on the ice
464 composition of carbonaceous chondrites from their magnetic mineralogy. *Earth Planet Sci*
465 *Lett* **576**, 117243 (2021).
- 466 44. Yabuta, H. *et al.* Macromolecular organic matter in samples of the asteroid (162173)
467 Ryugu. *Science (1979)* **379**, eabn9057 (2023).
- 468 45. Blum, J., Bischoff, D. & Gundlach, B. Formation of comets. *Universe* **8**, 381 (2022).
- 469 46. Busemann, H. *et al.* Interstellar chemistry recorded in organic matter from primitive
470 meteorites. *Science (1979)* **312**, 727–730 (2006).
- 471 47. Binkert, F. & Birnstiel, T. Carbon depletion in the early Solar system. *Mon Not R Astron*
472 *Soc* **520**, 2055–2080 (2023).

- 473 48. De Gregorio, B. T. *et al.* Isotopic and chemical variation of organic nanoglobules in
474 primitive meteorites. *Meteorit Planet Sci* **48**, 904–928 (2013).
- 475 49. Nakamura-Messenger, K., Messenger, S., Keller, L. P., Clemett, S. J. & Zolensky, M. E.
476 Organic globules in the Tagish Lake meteorite: Remnants of the protosolar disk. *Science*
477 (1979) **314**, 1439–1442 (2006).
- 478 50. Muñoz-Caro, G. M. *et al.* Comparison of UV and high-energy ion irradiation of methanol:
479 ammonia ice. *Astron Astrophys* **566**, A93 (2014).
- 480 51. Stammer, S. M. & Birnstiel, T. DustPy: a python package for dust evolution in
481 protoplanetary disks. *Astrophys J* **935**, 35 (2022).
- 482 52. Dullemond, C. P. *et al.* RADMC-3D: A multi-purpose radiative transfer tool. *Astrophysics*
483 *Source Code Library ascl*–1202 (2012).
- 484 53. Cruz-Diaz, G. A., Muñoz-Caro, G. M., Chen, Y.-J. & Yih, T.-S. Vacuum-UV
485 spectroscopy of interstellar ice analogs-I. Absorption cross-sections of polar-ice
486 molecules. *Astron Astrophys* **562**, A119 (2014).
- 487 54. Birnstiel, T., Fang, M. & Johansen, A. Dust evolution and the formation of planetesimals.
488 *Space Sci Rev* **205**, 41–75 (2016).

489

490

491 **Correspondence and requests for materials may be addressed to N.F.W. Ligterink.**

492

493 **Acknowledgments: The authors thank the three reviewers for their valuable input and**
494 **insights, which helped strengthen the paper.** The authors thank Eva G. Bøgelund and
495 Maria Drozdovskaya for insightful discussions.

496

497 **Author contributions:**

498 Conceptualization: NFWL, NvdM, AB, JTvS

499 Methodology: NFWL, PP,

500 Investigation: NFWL, PP, NvdM, JTvS, AB, CMOA, MEIR

501 Visualization: NFWL, NvdM, PP

502 Writing – original draft: NFWL, PP, NvdM, JTvS, AB, CMOA, MEIR

503 Writing – review & editing: NFWL, PP, NvdM, JTvS, AB, CMOA, MEIR

504

505 **Funding:** NFWL is supported by the Swiss National Science Foundation (SNSF) Ambizione
506 Grant 193453. PP acknowledge support from the UK Research and Innovation (UKRI)
507 under the UK government’s Horizon Europe funding guarantee from ERC. MEIR is
508 supported by the SNSF Ambizione grant 193331.

509

510 **Competing interests:** Authors declare that they have no competing interests.

511

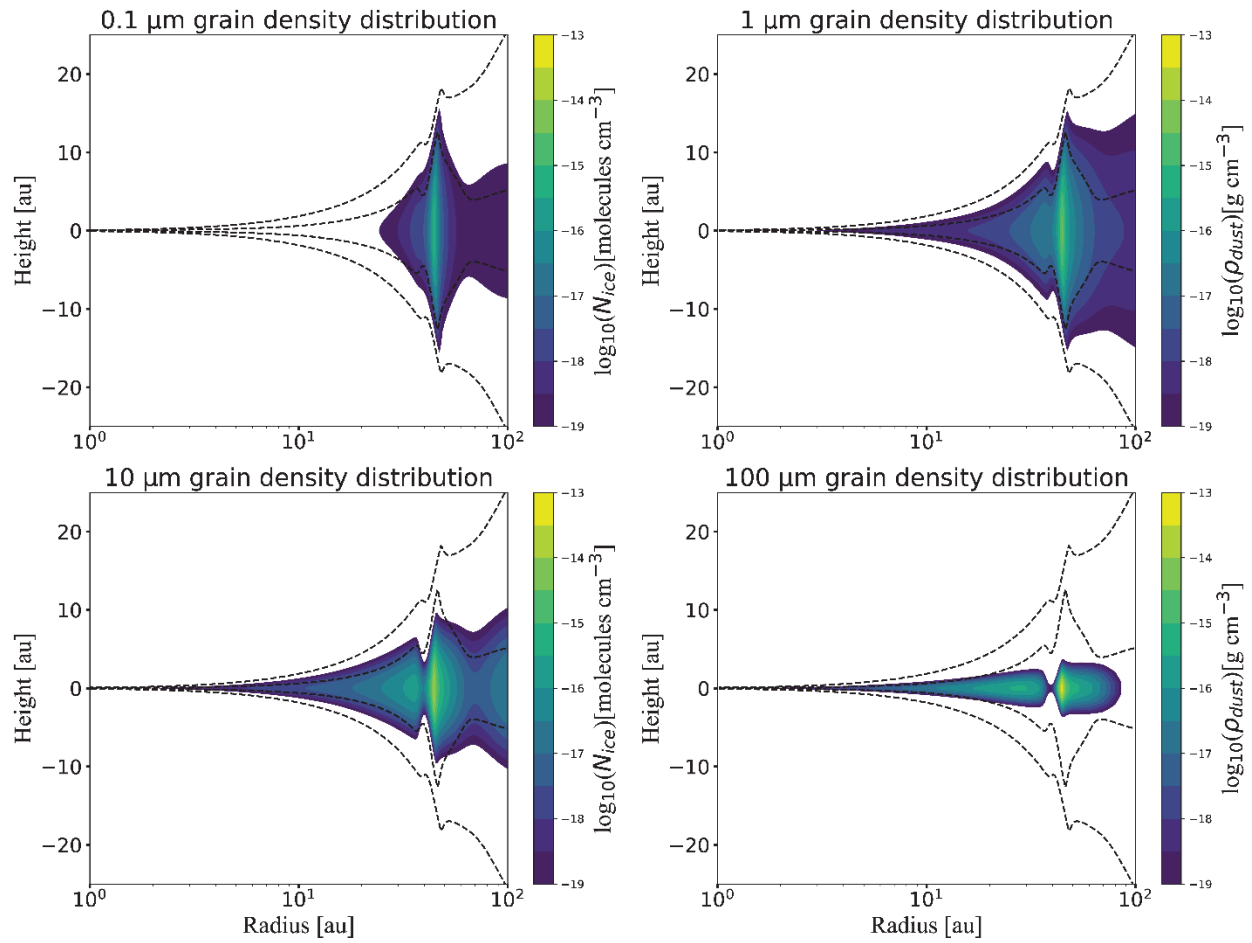
512 **Data and materials availability:** Data and code are available upon request.

513

514 **Extended Data**

515

516



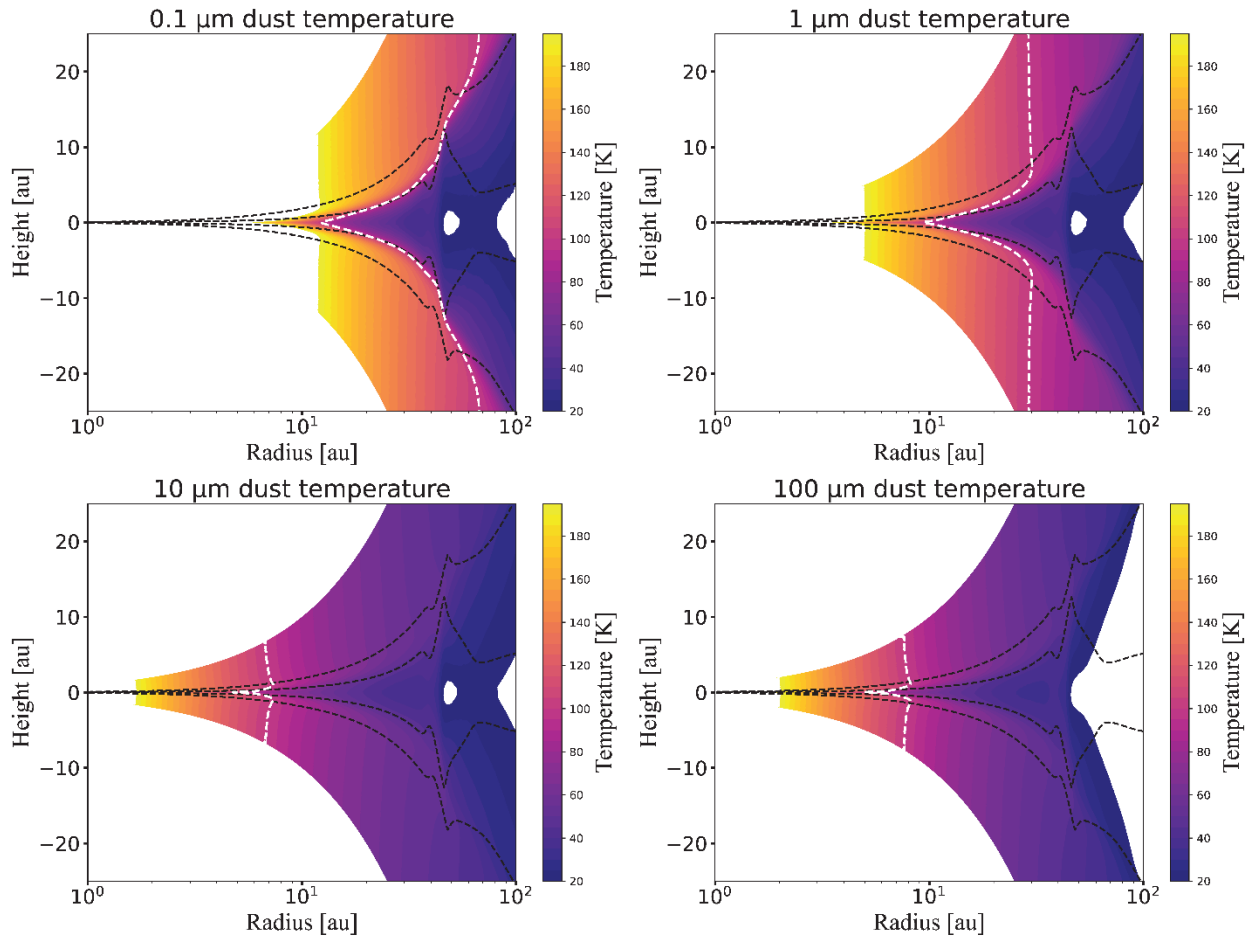
517

518 **Fig. S1.** Dust density distribution for grains of 0.1, 1, 10, and 100 μm in size. The black dashed
519 lines indicate the total dust density distribution contours at 10^{-19} and 10^{-16} g cm^{-3} .

520

521

522



523

524

525

526

527

528

529

Fig. S2. Dust temperatures for grains of 0.1, 1, 10, and 100 μm in size. The white dashed lines indicate the 100 K temperature contour, while the black dashed lines indicate the total dust density distribution contours at 10^{-19} and 10^{-16} g cm^{-3} .